



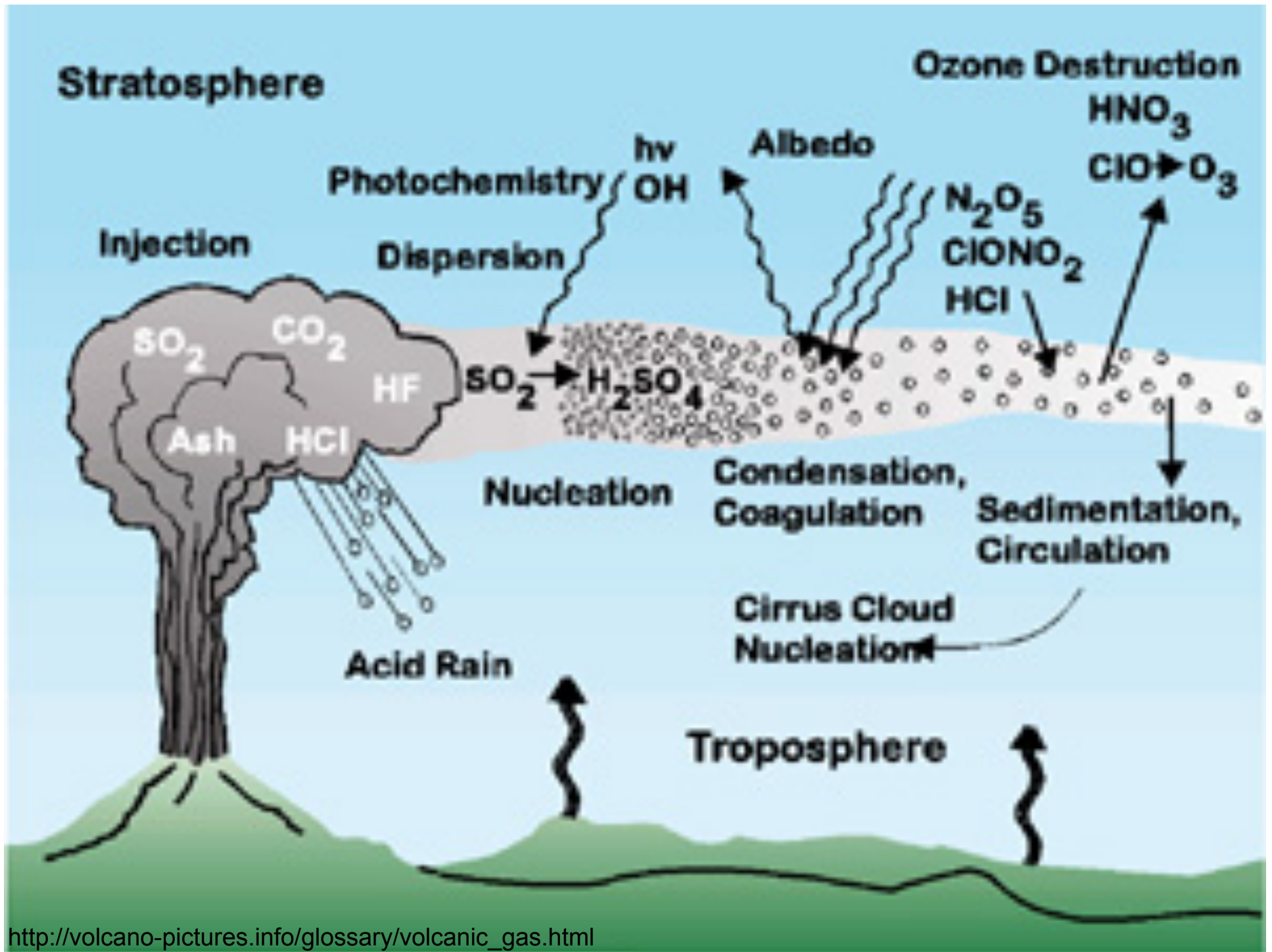
# Chemistry of SO<sub>2</sub> in tropospheric volcanic plumes

*by*

*Dr. Lizzette A. Rodríguez Iglesias*

**Department of Geology  
University of Puerto Rico  
Mayagüez Campus**

*Photo: L. Rodriguez*



# Introduction

- $\text{SO}_2$  emission rates are commonly measured using ground-based spectroscopic techniques, a few km downwind from the vent and km from the plume
  - This allows time for  $\text{SO}_2$  to interact with other volcanogenic gases, particles and droplets of volcanogenic and/or meteoric origin, as well as atmospheric gases and aerosols, and for the possibility of attenuation of the  $\text{SO}_2$  signal by airlight or scattering.
  - Measured  $\text{SO}_2$  flux will not usually represent the at-source  $\text{SO}_2$  emission rate



# Cont. Introduction

- An important  $\text{SO}_2$  depletion process is the conversion of  $\text{SO}_2$  to  $\text{SO}_4^{-2}$  (sulfate aerosols)
  - $\text{SO}_2$  loss rates ranging from  $10^{-7}$  to  $10^{-3} \text{ s}^{-1}$  have been estimated for tropospheric volcanic plumes at various altitudes.
- Reactions  $\text{SO}_2$  can undergo, leading to formation of particulate sulfate: gas-phase homogeneous (slower: days to weeks), aqueous-phase (hours), heterogeneous reactions on the surface of solids.



# Cont. Introduction

- **Local meteorology** affects the fate of tropospheric plumes both directly, through dispersion and transport downwind, and indirectly, through factors such as humidity, T, amount of sunlight reaching the plume, cloud cover, fog, and precipitation.
- **Other factors affecting reactions:** aerosol concentration and pH, availability of oxidants (e.g. OH, O<sub>3</sub>, H<sub>2</sub>O<sub>2</sub>)

# Methodology

- SO<sub>2</sub> fluxes are used to obtain loss rates

$$k, \quad \phi_{t1} = \phi_{t2} e^{k1(t2-t1)}$$

where  $\phi_t$  represents an SO<sub>2</sub> flux at a given time t

*Photo: L. Rodriguez*

# Why?

- **Uncertainties in near-source plume chemistry can complicate interpretations of volcanic activity and hazards, petrology, global emission rates\* and climatic effects of emissions.**
- **Consideration of variable SO<sub>2</sub> depletion rates in volcanic plumes could contribute to future modeling of global S sources and distributions, as well as relative contributions.**



# Why?

- Volcanic sulfate aerosols:
  - injected into the free troposphere, where removal processes are slow
  - uncertainties in their contribution are important for climate change studies
  - cool the climate due to backscattering of sunlight and through an increase in cloud reflectivity and residence time
  - can absorb outgoing LW terrestrial radiation
- 18-40% of the global tropospheric sulfate burden is volcanogenic

# Previous Work

★ Eatough et al., 1994

✧ Anthropogenic emissions of SO<sub>2</sub> and chemistry of the conversion

★ Thornton et al., 1996

✧ Measurements in remote marine areas

★ Oppenheimer et al., 1998

✧ SHV measurements (1996)

★ Horrocks et al., 2003

✧ FTIR Masaya – negligible loss

★ McGonigle et al., 2004

✧ Depletion rates at Masaya – negligible loss

★ Nadeau and Williams-Jones, 2009

✧ Depletion rates at Masaya – 33-50% less @15 km away: conversion negligible, probably caused by dilution of plume by greater wind speeds at different altitudes. Apparent loss.

# SO<sub>2</sub> loss rates in the plume emitted by Soufrière Hills volcano, Montserrat



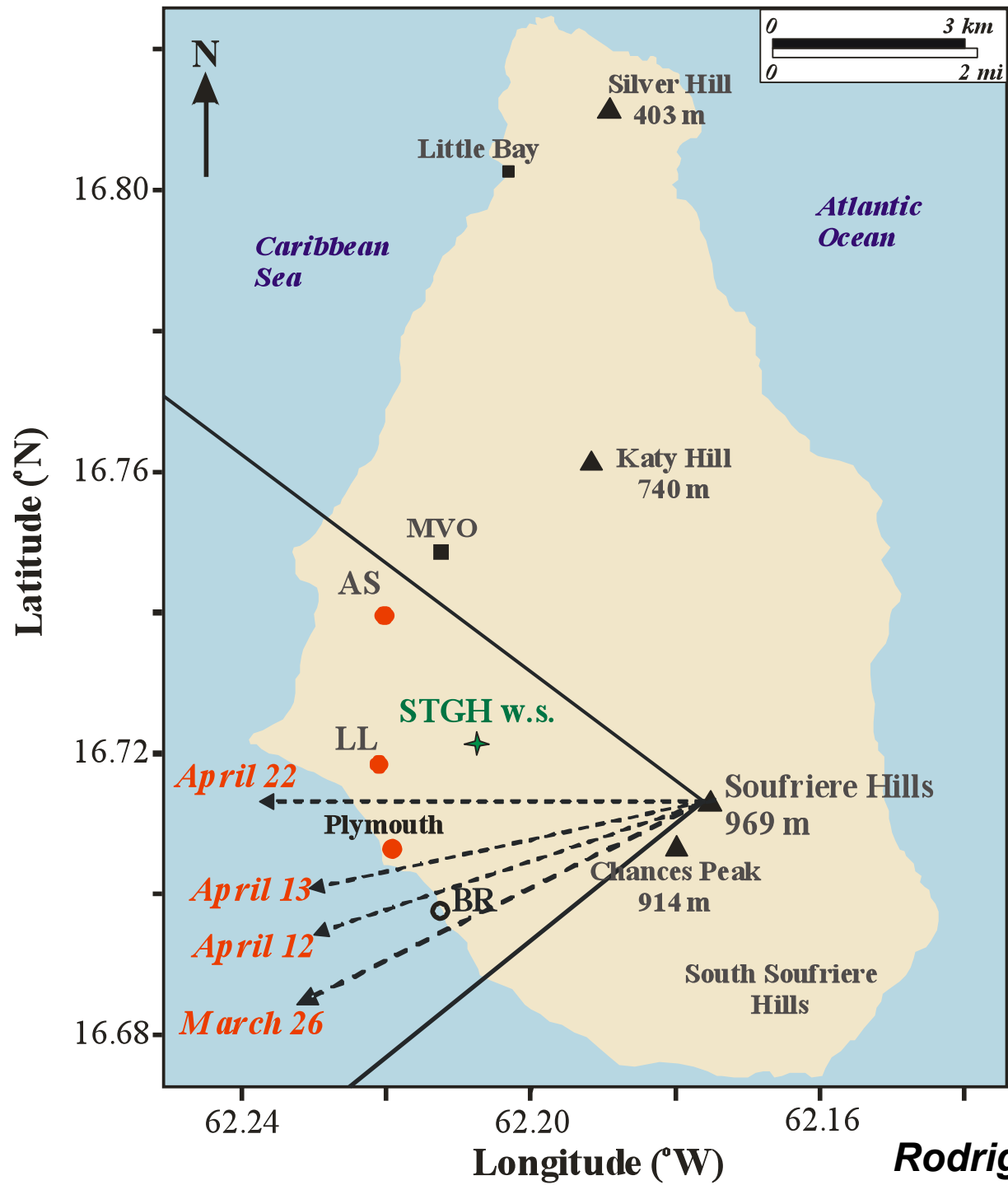
Rodríguez, L.A. et al., “SO<sub>2</sub> loss rates in the plume emitted by Soufrière Hills volcano, Montserrat”, *JVGR*, 173 (1-2): 135-147 (DOI 10.1016/j.jvolgeores.2008.01.003).

*Photo: L. Rodriguez*



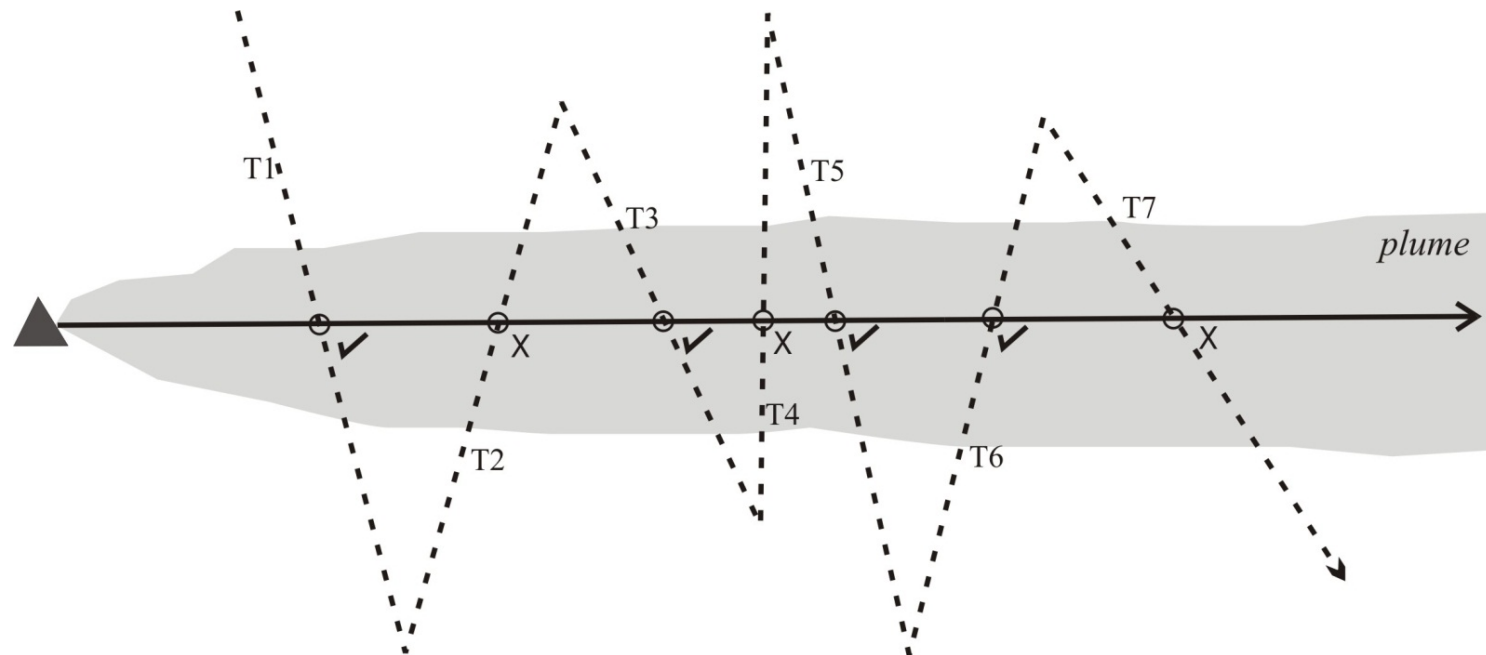
# Objectives

- Quantify SO<sub>2</sub> depletion rates in volcanic plumes injected into the boundary layer, from tropical low altitude volcanoes in a humid environment (typical of ~20% of active volcanoes worldwide), using ground-based remote sensing techniques
- SO<sub>2</sub> fluxes were measured near to the eruptive vent and at various distances downwind of the Soufrière Hills volcano (SHV)



*Rodriguez et al., 2008*

# Plume tracking to follow a plume portion downwind



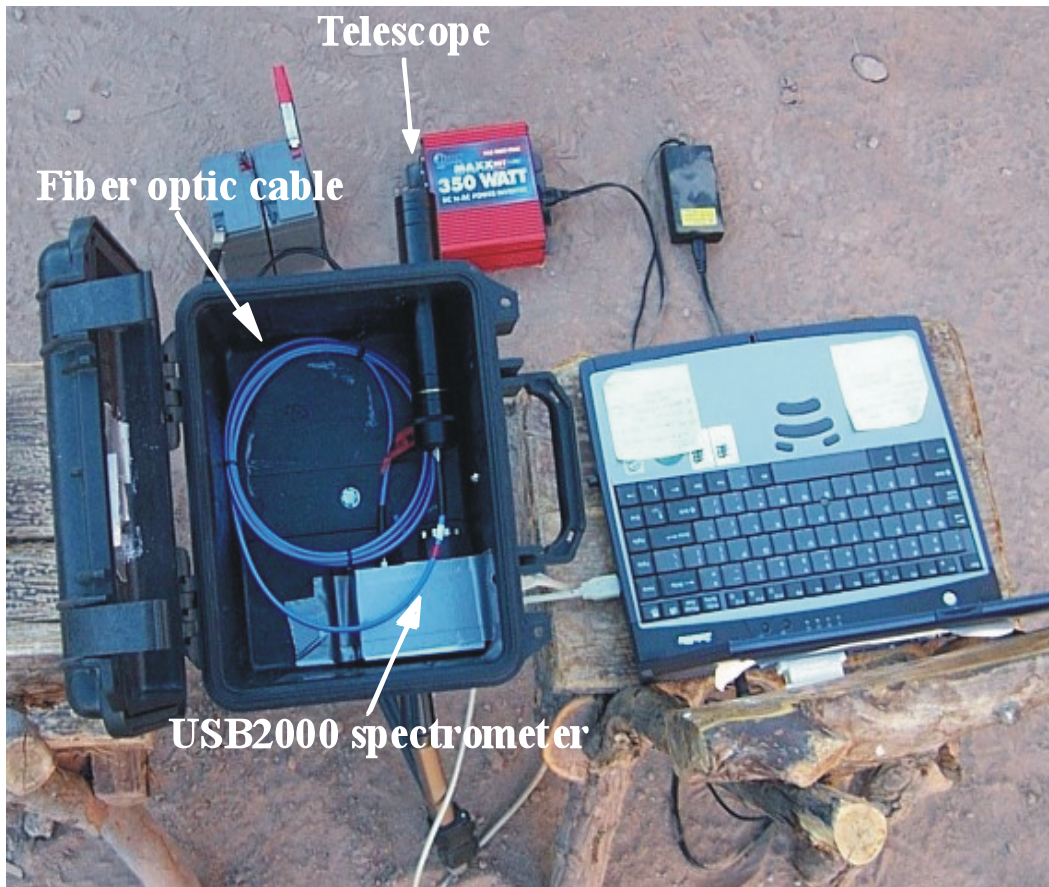
- ▲ Vent
- Plume azimuth
- - → Helicopter track
- Intersections between helicopter track and plume azimuth
- T<sub>x</sub> Traverse number
- ✓ SO<sub>2</sub> fluxes used to calculate loss rates for a plume portion traveling from t1
- X SO<sub>2</sub> fluxes that did not follow criteria and were not used for loss rate calculation

**Rodriguez  
et al., 2008**



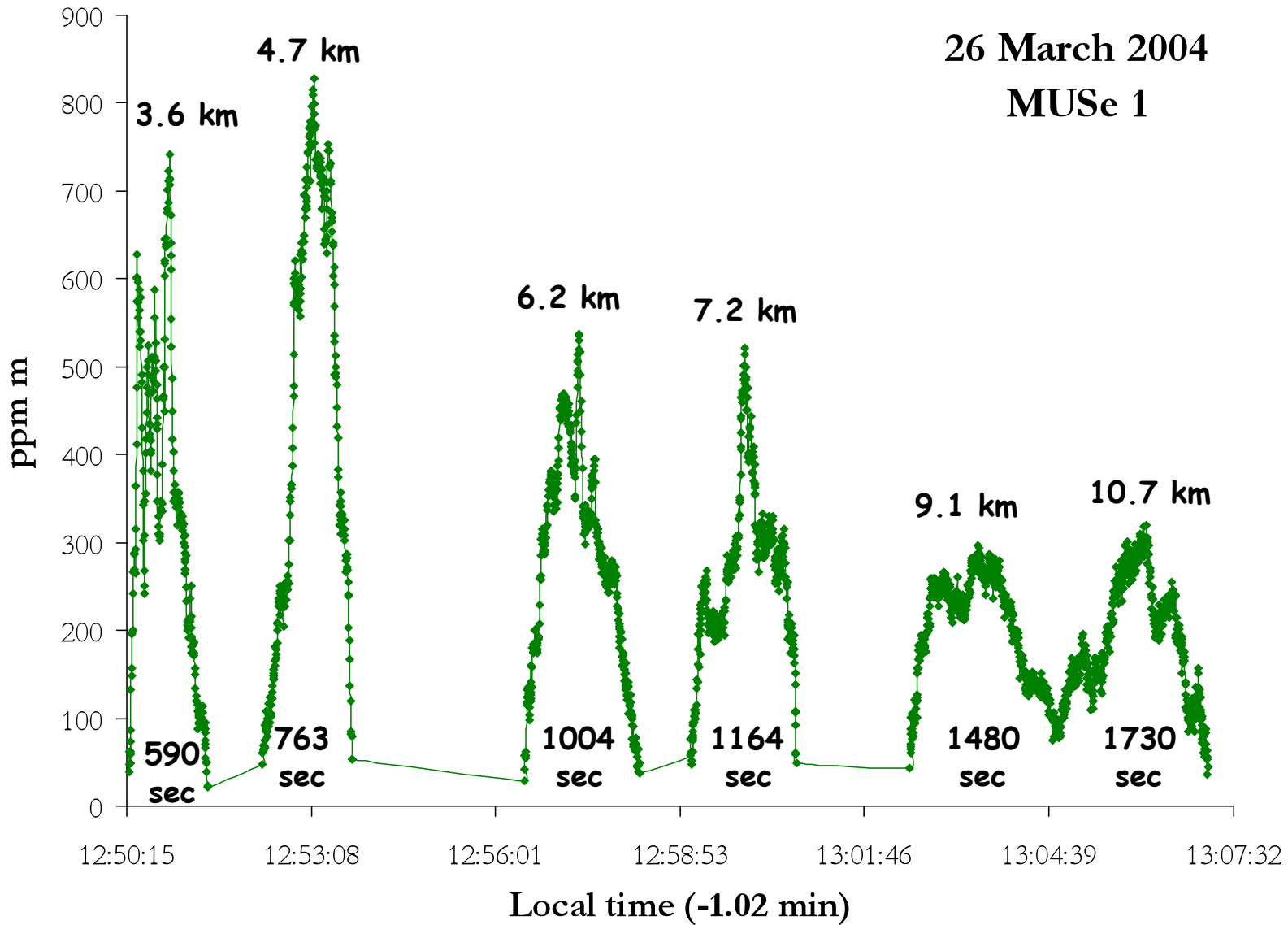
**Equipment: (a) Mini-UV spectrometer (MUSE), optical assembly and laptop computer used for measurements.  
(b) Telescope setup**

**(a)**



**(b)**



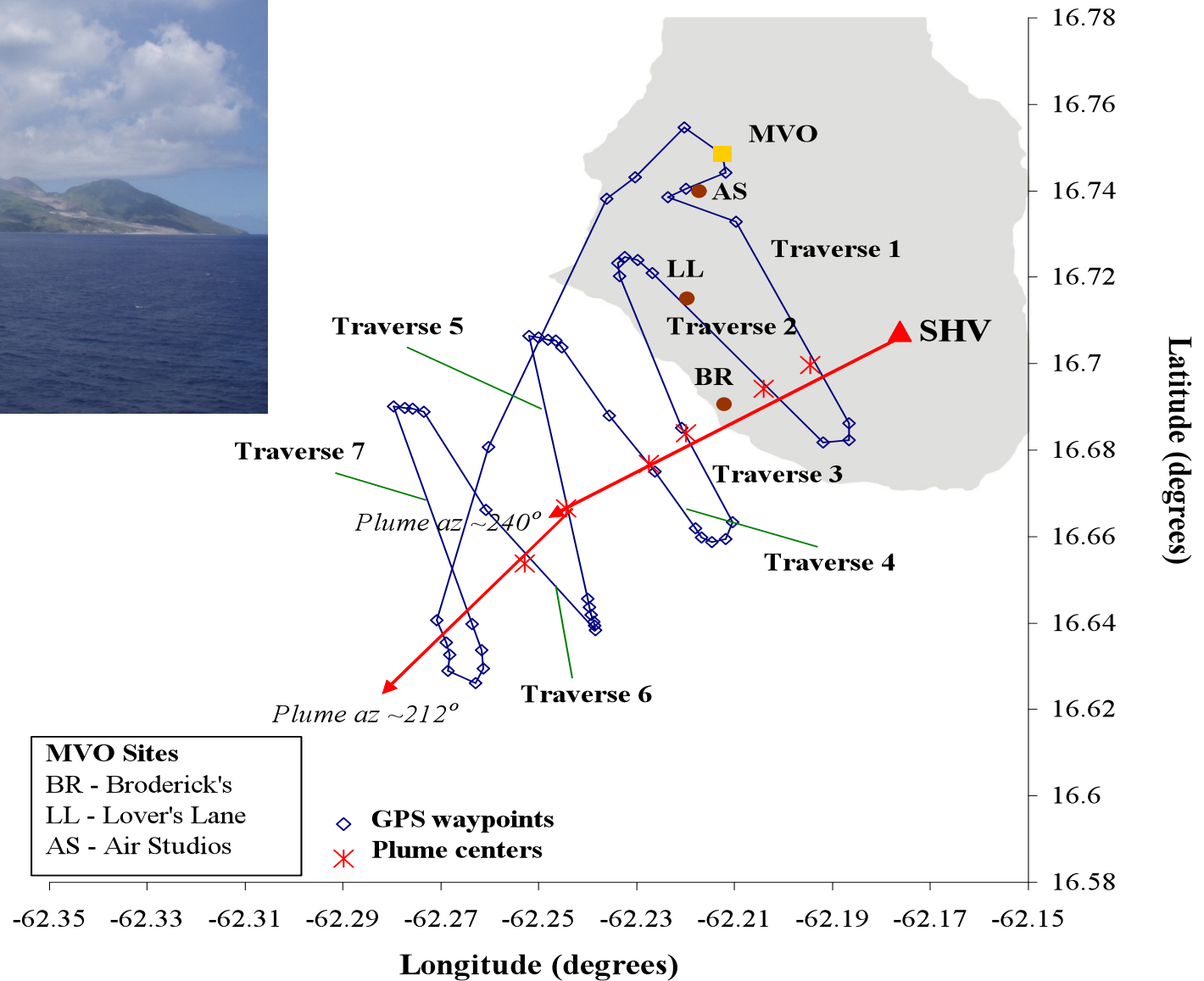


***By: L. Rodriguez***

# 26 March 2004



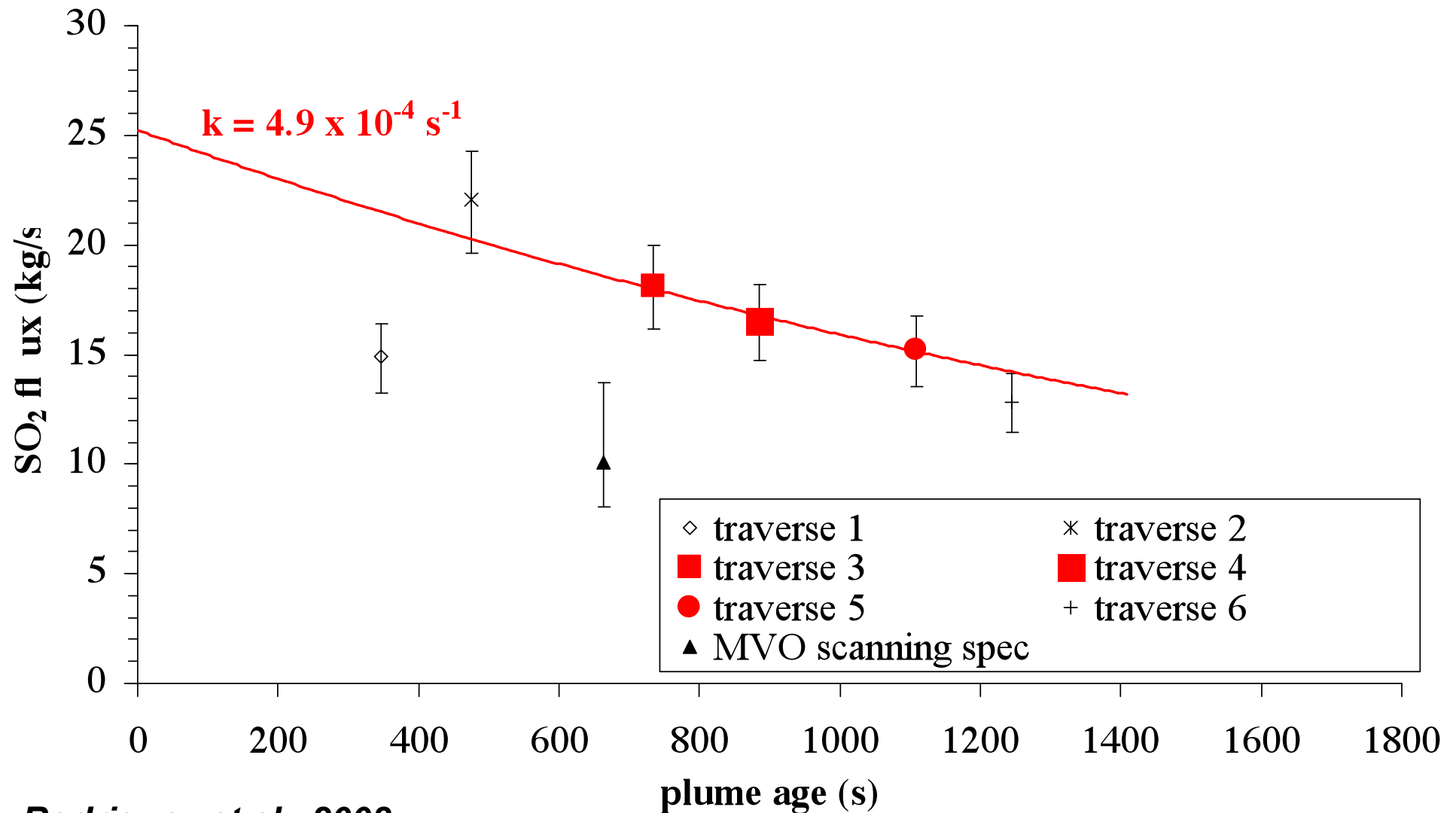
\*Plume centers are 2.1 to 11.9 km from the vent (coastline at plume az is ~4.2 km from vent).



Rodriguez  
et al., 2008



# SO<sub>2</sub> flux vs. plume age (a best-fit exponential curve yields the flux at time=0 s: 25 kg/s (2160 t/d))



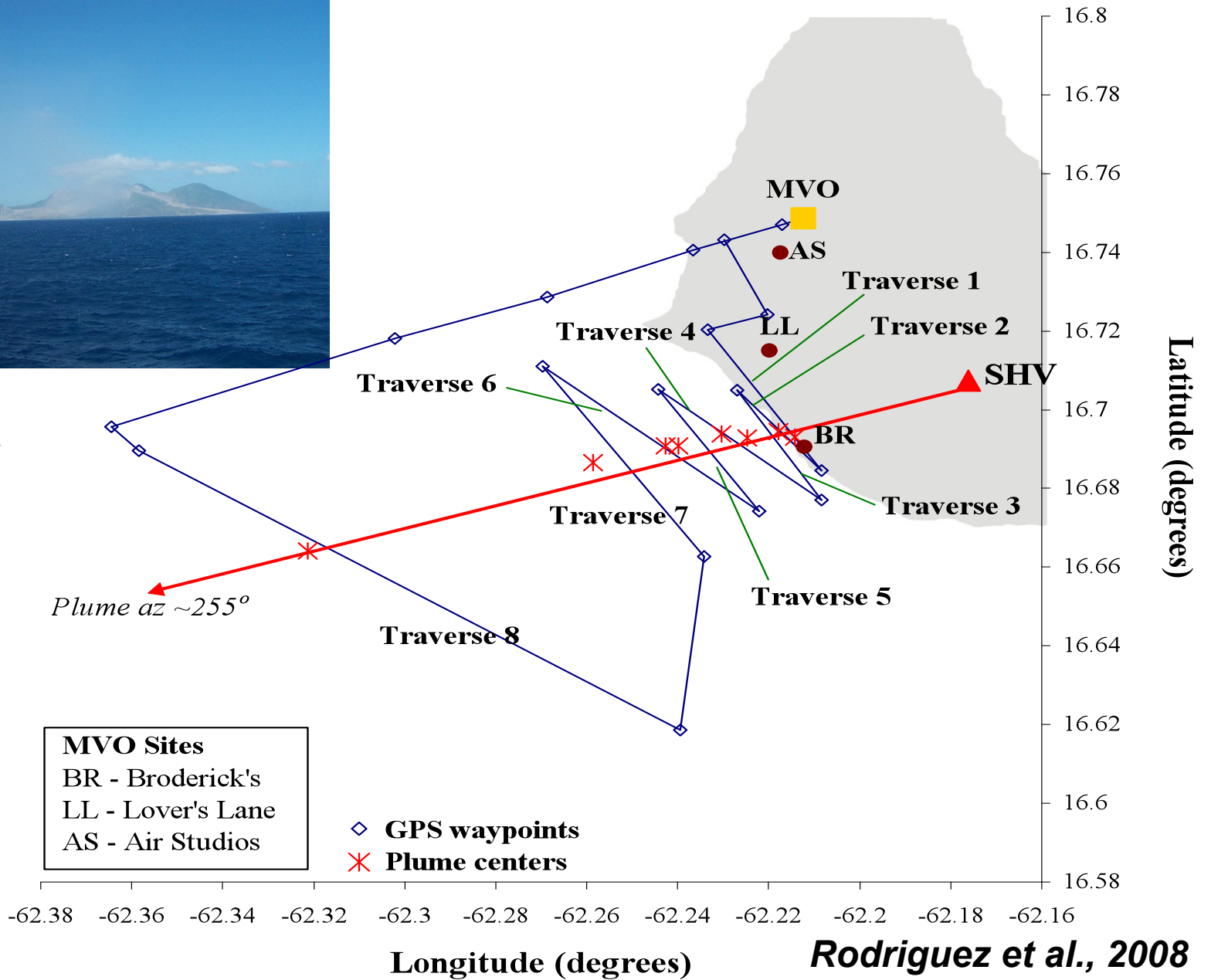
Rodriguez et al., 2008

Error bars are -11% +10%. BR's ave emission rate - error bar is -20% +36%.

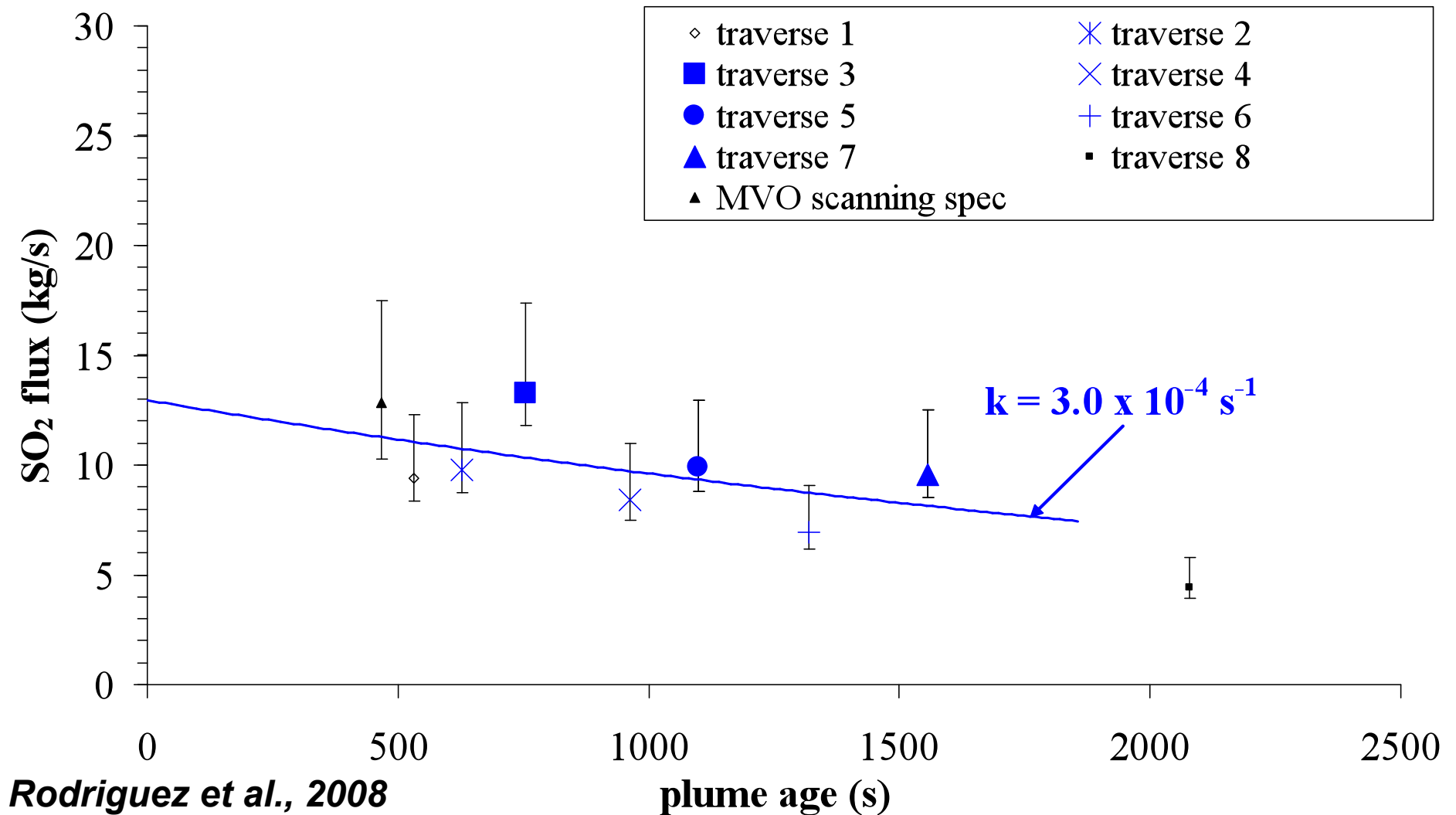
# 13 April 2004



\* Plume centers are 4.3 to 16.2 km from the vent (coastline at plume az is ~4.4 km from vent).



# SO<sub>2</sub> flux vs. plume age (a best-fit exponential curve yields the flux at time=0 s: 13 kg/s (1120 t/d))



Error bars are -11% +31%. BR's ave emission rate - error bar is -20% +36%.

# Discussion

- SO<sub>2</sub> fluxes measured at progressively larger distances (and hence plume ages) downwind decrease at a predictable rate.
  - Calculated loss rates ranged from  $3.0 \times 10^{-4} \text{ s}^{-1}$  to  $1.4 \times 10^{-3} \text{ s}^{-1}$  (e-folding times: 0.9-0.2 hrs): ascribed mainly to heterogeneous removal of SO<sub>2</sub>.
- Extrapolation of data back to  $t = 0 \text{ s}$  gives an average of the at-source emission rate, based on an exponential decay of SO<sub>2</sub>.
  - Results reported by MVO were significantly lower than ours, which take into account SO<sub>2</sub> loss. Further validation and comparison of the techniques would be of value.

# Cont. Discussion

- Similar study in 1996 (Oppenheimer et al., 1998): loss rates  $\sim 10^{-3} \text{ s}^{-1}$ , an order of magnitude faster than ours, which included more traverses over a longer period.
  - Conditions differed in that we measured ash-free plumes during the dry season, while Oppenheimer et al. (1998) measured ash plumes during the peak of the rainy season (greater concentration of available condensed atmospheric water).
  - Criteria used to choose the traverses is different



# Conclusions

- These differences are carried forward to models, and introduce an error.
- Contributions of volcanoes to the global  $\text{SO}_2$  budget are underestimated.

Apparent decreases in SO<sub>2</sub> flux  
in the plume emitted by  
Láscar volcano, Chile

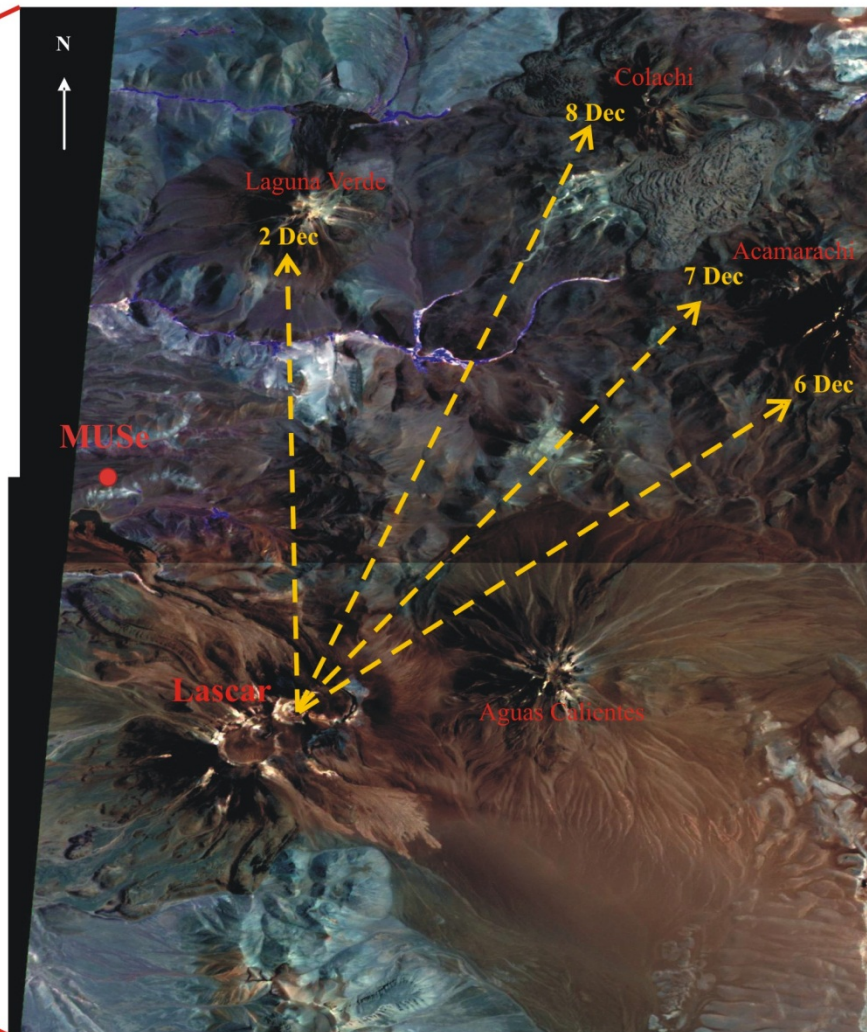
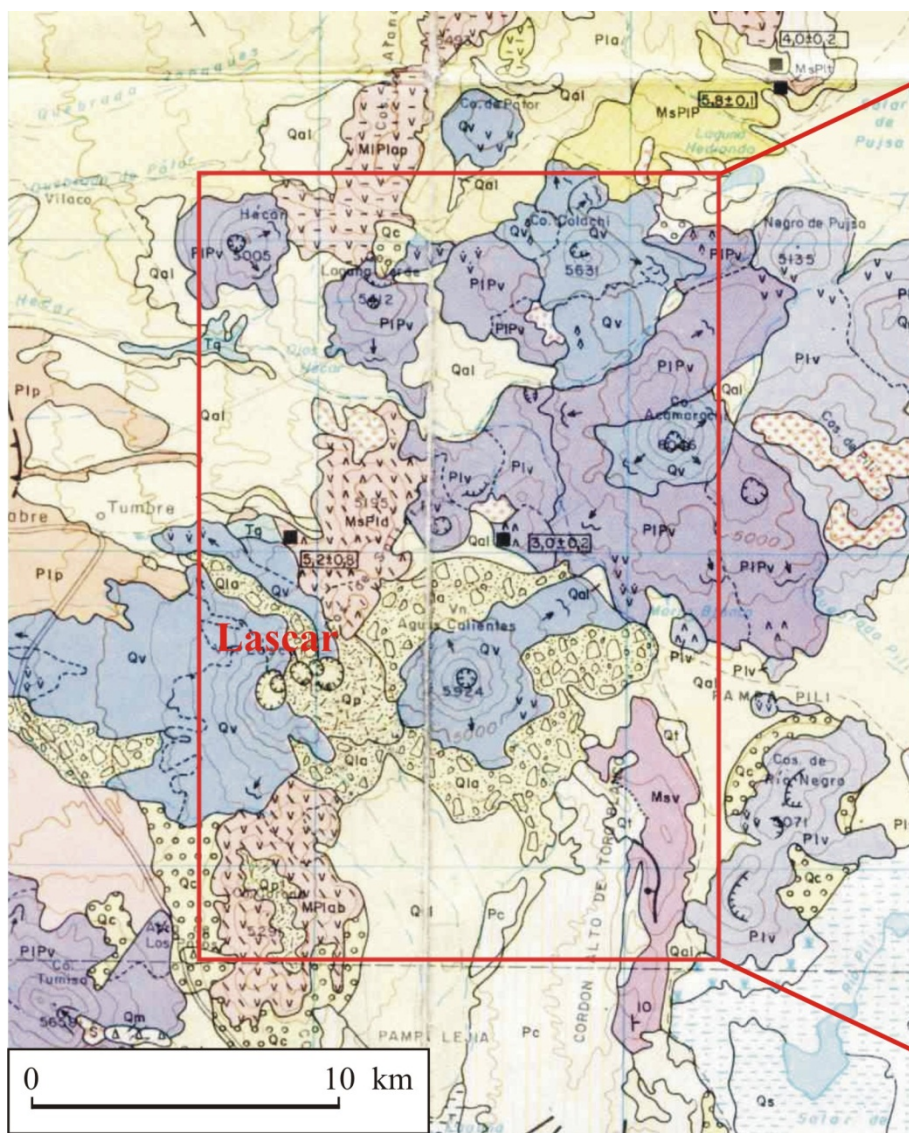


Rodríguez, L.A. et al., in prep.

*Photo: L. Rodriguez*



# Location of Láscar, approximate position of MUSE during measurements, and plume azimuths



ASTER image (1-2-3 visible composite)

By: L. Rodriguez

Toconao quadrangle (1:250,000)



# MUSe measurements



December 7, 2004. Plume traveling to the NE.

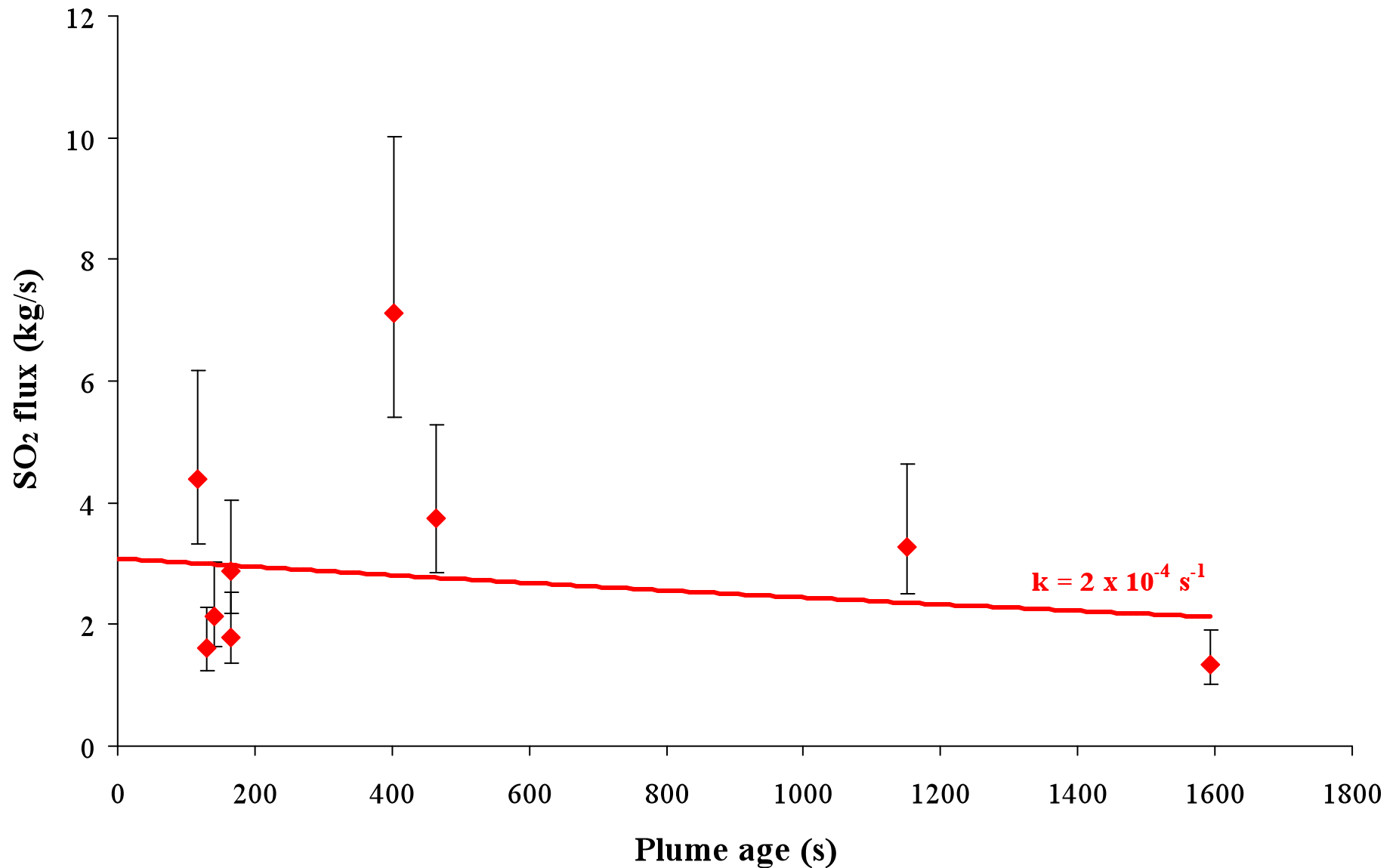


2 MUSe's, one measuring upwind (close to vent) and the other measuring downwind

*Photos: L. Rodriguez*



# SO<sub>2</sub> flux vs. plume age (7 Dec) (best-fit exponential curve yields flux at time=0 s: 3 kg/s (260 t/d))



# Discussion

- SO<sub>2</sub> fluxes averaged  $\sim 200 \text{ t d}^{-1}$ , an order of magnitude lower than previous measurements of  $2300 \text{ t d}^{-1}$  on Jan 2003 (Mather et al., 2004)
- Calculated loss rates ranged from  $2.0 \times 10^{-4} \text{ s}^{-1}$  to  $7 \times 10^{-4} \text{ s}^{-1}$  (e-folding times: 1.4-0.4 hrs), which are subject to large errors.
  - Most accurate estimate:  $2.0 \times 10^{-4} \text{ s}^{-1}$  (Dec 7).

# Cont. Discussion

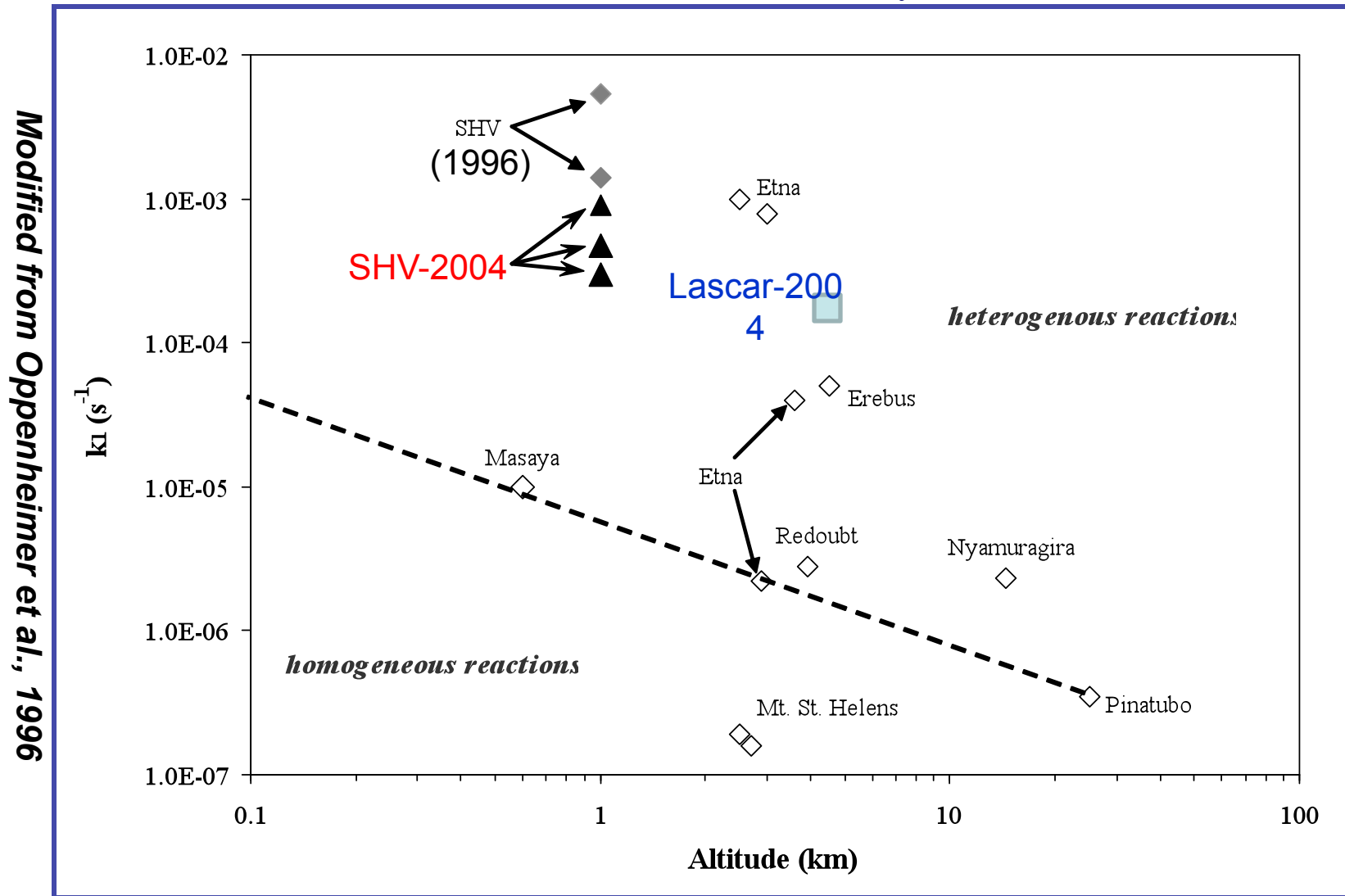
- Apparent loss of SO<sub>2</sub> because of the errors in the calculation of fluxes:
  - UV scattering causing attenuation as distance between instrument and plume increases – causes underestimates in the flux measurements
  - Errors in the plume azimuth
- These will produce an apparent faster loss of SO<sub>2</sub> in the plume

# Conclusions

- SO<sub>2</sub> loss rates obtained are a combination of depletion of SO<sub>2</sub> by heterogeneous reactions, of the effects of errors in the plume azimuth calculation, and of the long distances between the instruments and the plume
  - Latter effect due to UV scattering, which can produce decreases in the flux and consequently make the loss rates appear faster, without any removal mechanism acting on the plume.
  - Residence times are longer than calculated



# SO<sub>2</sub> depletion rates as a function of altitude of the plume



Dashed line indicates the approximate transition between homogeneous and heterogeneous reactions (based on Eatough et al., 1994 and Thornton et al., 1996).